

The TRANSMED Atlas

The Mediterranean Region from Crust to Mantle

Geological and Geophysical Structure of the
Mediterranean and the Surrounding Areas

A publication of the Mediterranean Consortium
for the 32nd International Geological Congress

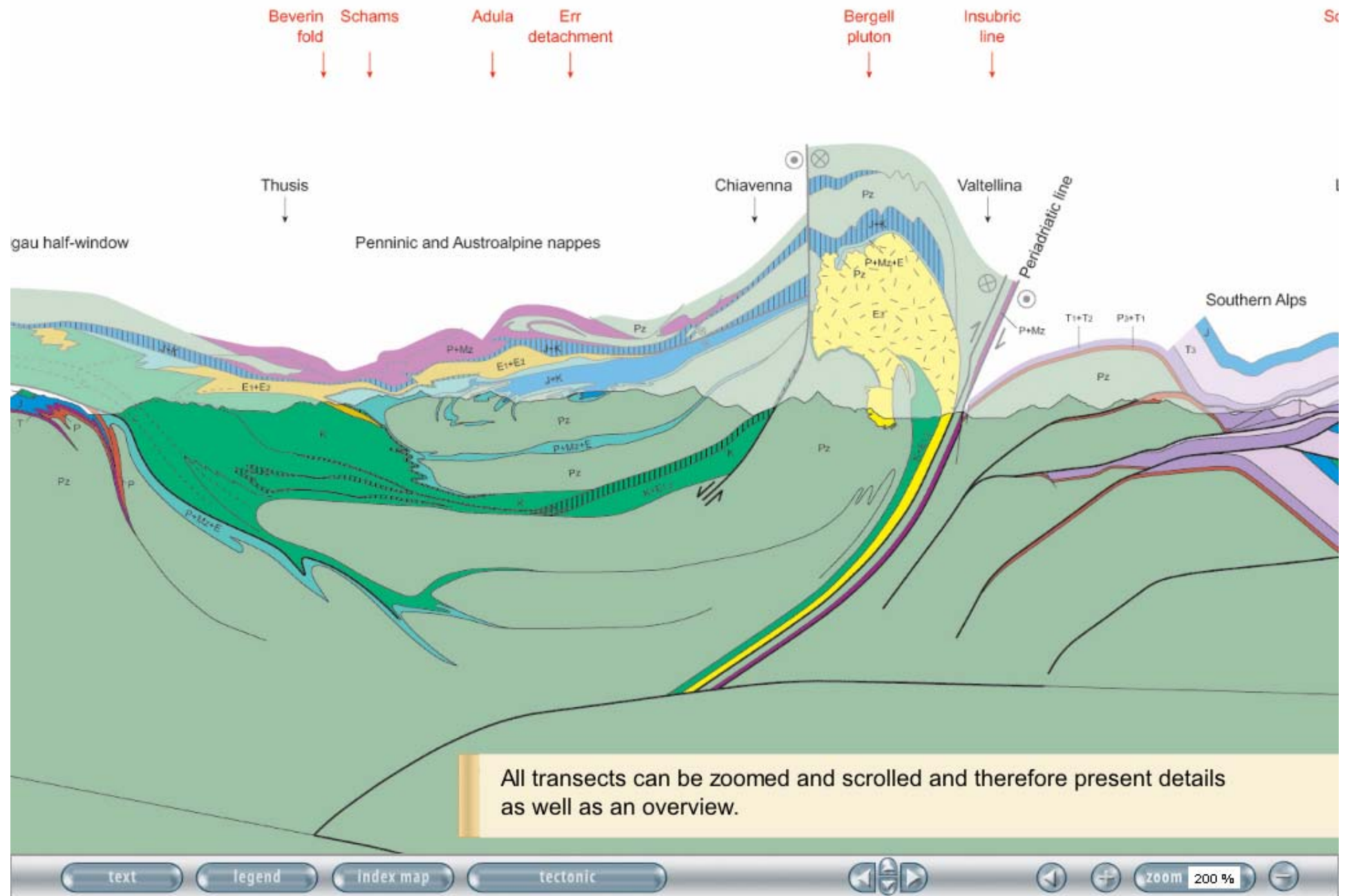


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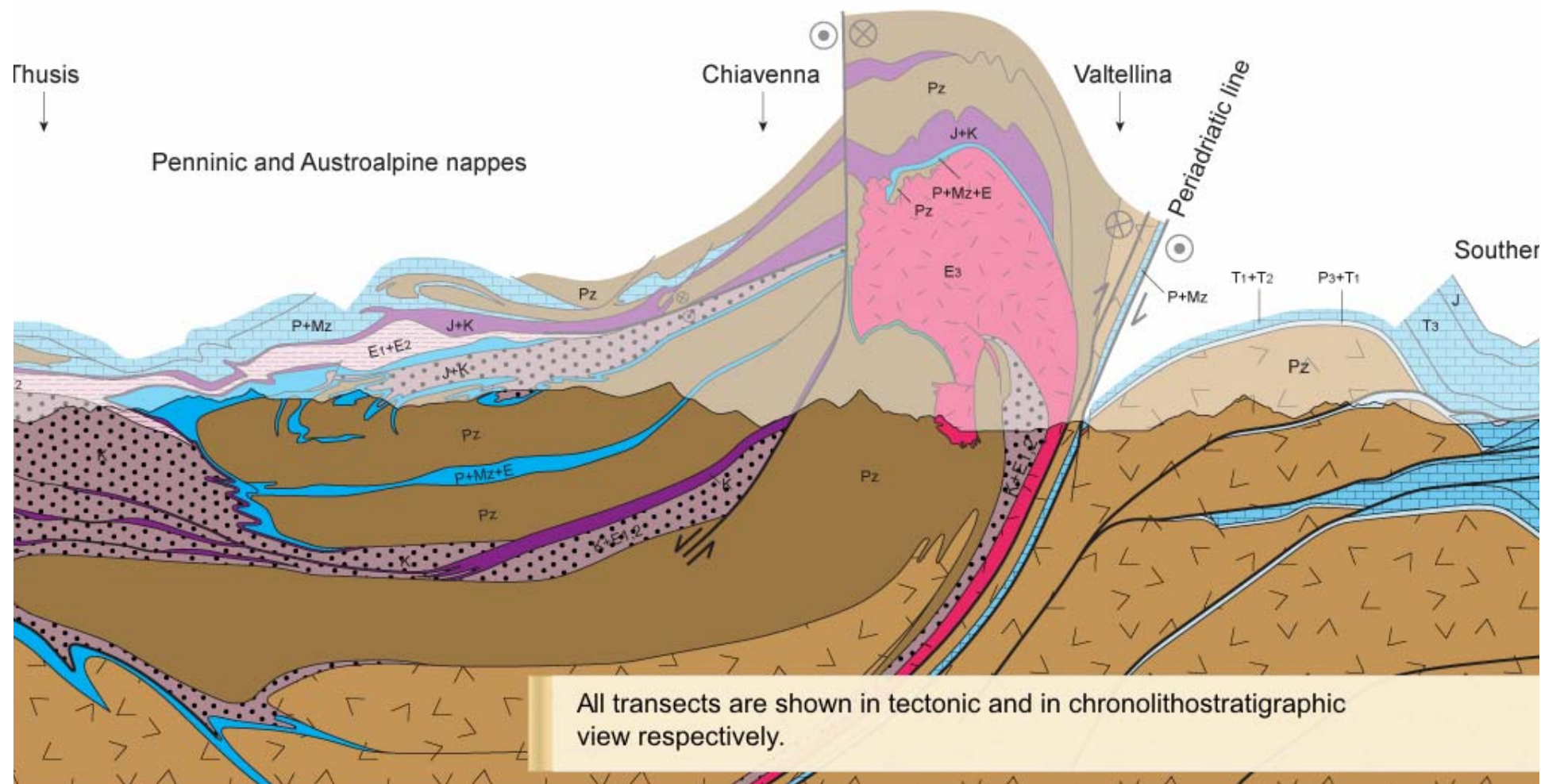


The main menu gives an overview over all transects and direct access to all detail information on each transect.



Beverin fold
Schams
Adula
Err detachment

Bergell pluton
Insubric line



All transects are shown in tectonic and in chronolithostratigraphic view respectively.

text

legend

index map

chronolithostratigraphic

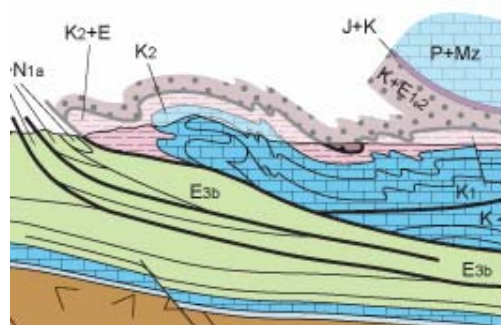


zoom 200 %

Saentis thrust



Helvetic nappes



Additional Information



Austroalpine nappes (Tilisuna)

Gently N-dipping basal thrust of the Austroalpine Silvretta nappe, above the upper Penninic Arosa unit and the middle Penninic Sulzfluh unit (Jurassic platform carbonates) near Tilisuna hut in the Austrian-Swiss border area (Montafon).

Additional information on some features of the transects can be accessed by links.

text

legend

index map



chronolithostratigraphic







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Legend






Tectonic basins

-  foreland basin
-  piggy-back basin

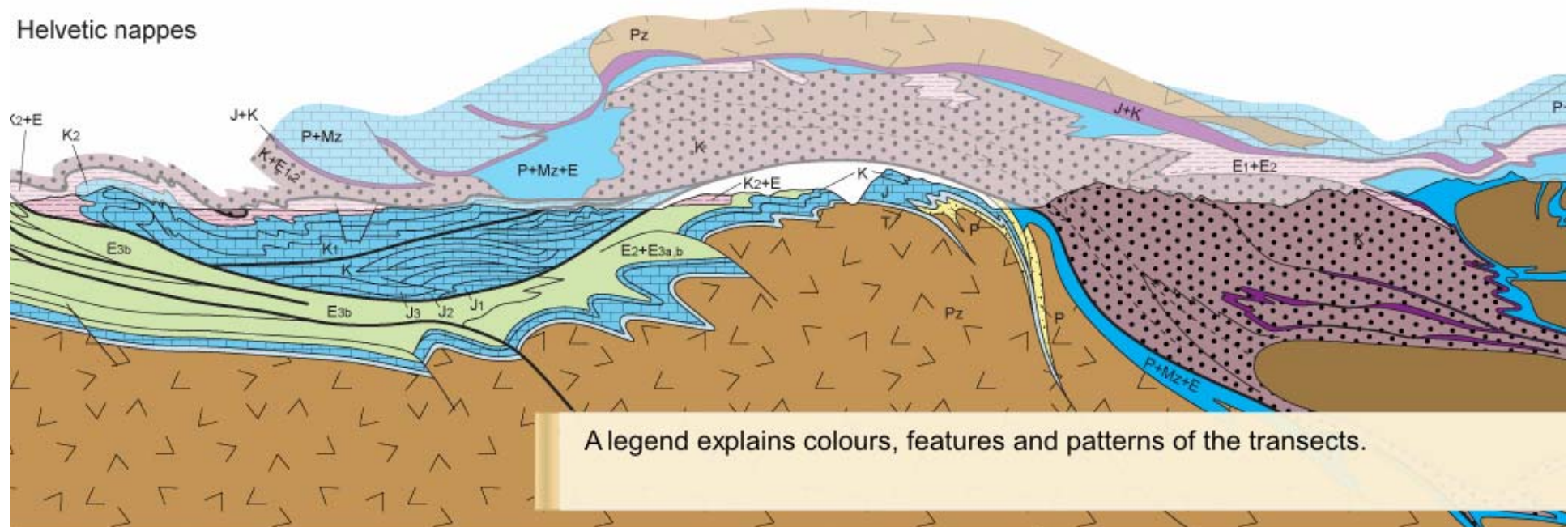
Shelf/slope/rise

-  shelf neritic & paralic clastics
-  shelf carbonates
-  shelf basins
-  slope/rise undifferentiated

Continental crust

-  indeterminate sediments (Paleozoic)
-  nonmarine clastics (late Paleozoic rifts)
-  rifted continental crust
-  upper continental crust
-  lower continental crust

Helvetic nappes



A legend explains colours, features and patterns of the transects.

text

legend

index map

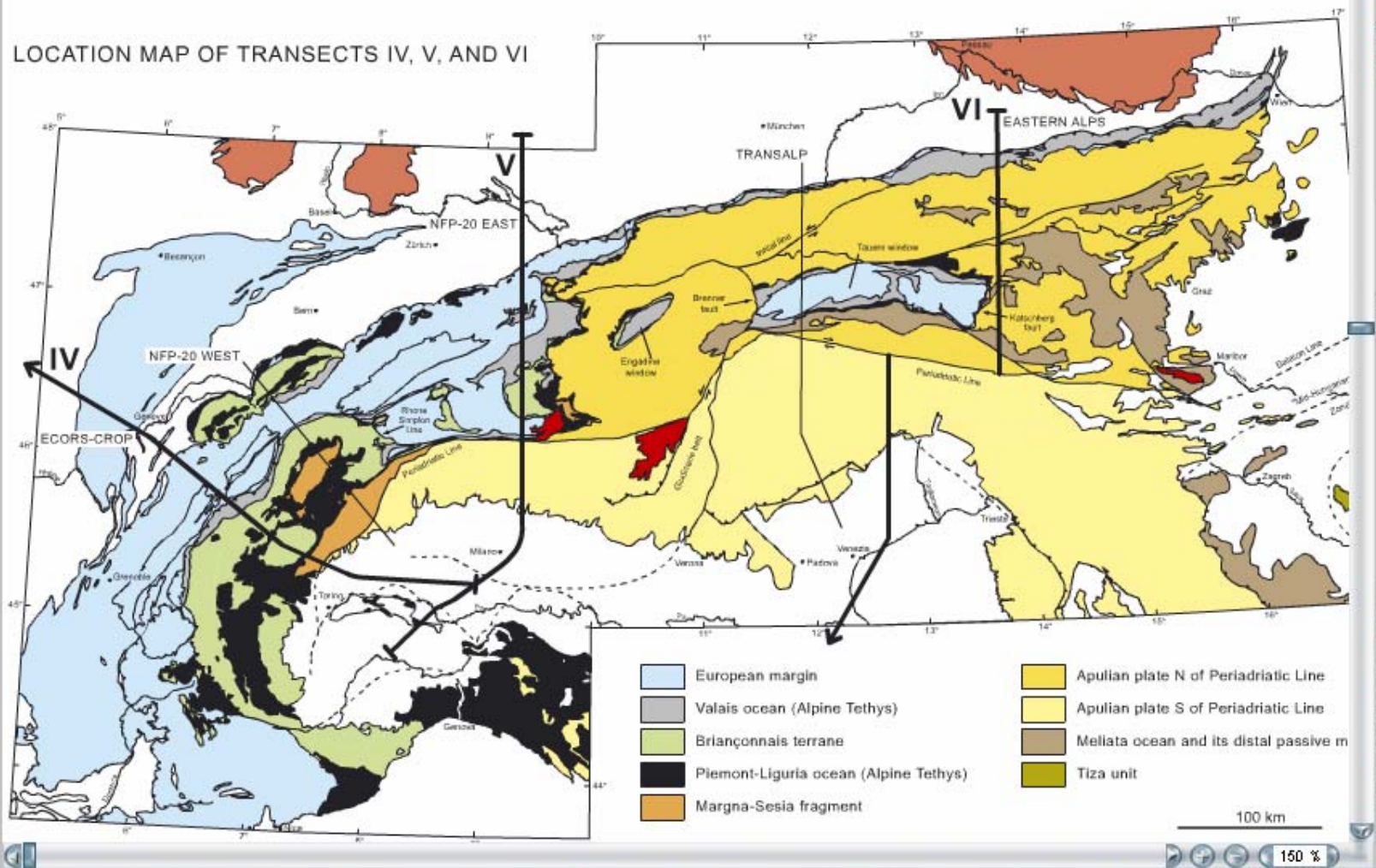
chronolithostratigraphic



zoom 200 %

Index Map

LOCATION MAP OF TRANSECTS IV, V, AND VI



erin fold
Schams

Penninic and /

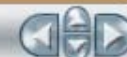
In every transect an index map shows the exact location of the transect.

text

legend

index map

chronolithostratigraphic



zoom 200 %

Transect IV-VI

TRANSMED Transects IV, V and VI: Three lithospheric transects across the Alps and their forelands

Stefan M. Schmid, Bernhard Fügenschuh, Eduard Kissling, Ralf Schuster

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 - [4.1. TRANSMED transect IV \(ECORS-CROP\)](#)
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1. General Introduction

The European Alps, located in south-central Europe, record the closure of several ocean basins located in the Mediterranean domain during the Late Cretaceous and Cenozoic convergence of the African (or Apulian) and European plates (e.g. Trümpy 1960; Frisch 1979; Haas et al. 1995; Stampfli et al. 2001a,b). In recent years it has become increasingly evident that the oceanic and continental paleogeographical realms from which the Alpine tectonic units derive were arranged in a rather non-cylindrical fashion. This led to important along-strike changes in the overall architecture of the Alps (Fig. [IV-VI.1](#)), also reflected, for example, in the deep structure of the Alps (e.g. Pfiffner et al. 1997; Schmid and Kissling 2000) or in the different age of the main metamorphic events: Tertiary in the Western Alps, Cretaceous in the Austroalpine units of the Eastern Alps (e.g. Gebauer, 1999; Thöni, 1999).

As discussed in this paper (see also

The results and interpretation of the transect-data can be accessed directly from the panorama.

Transect IV-VI

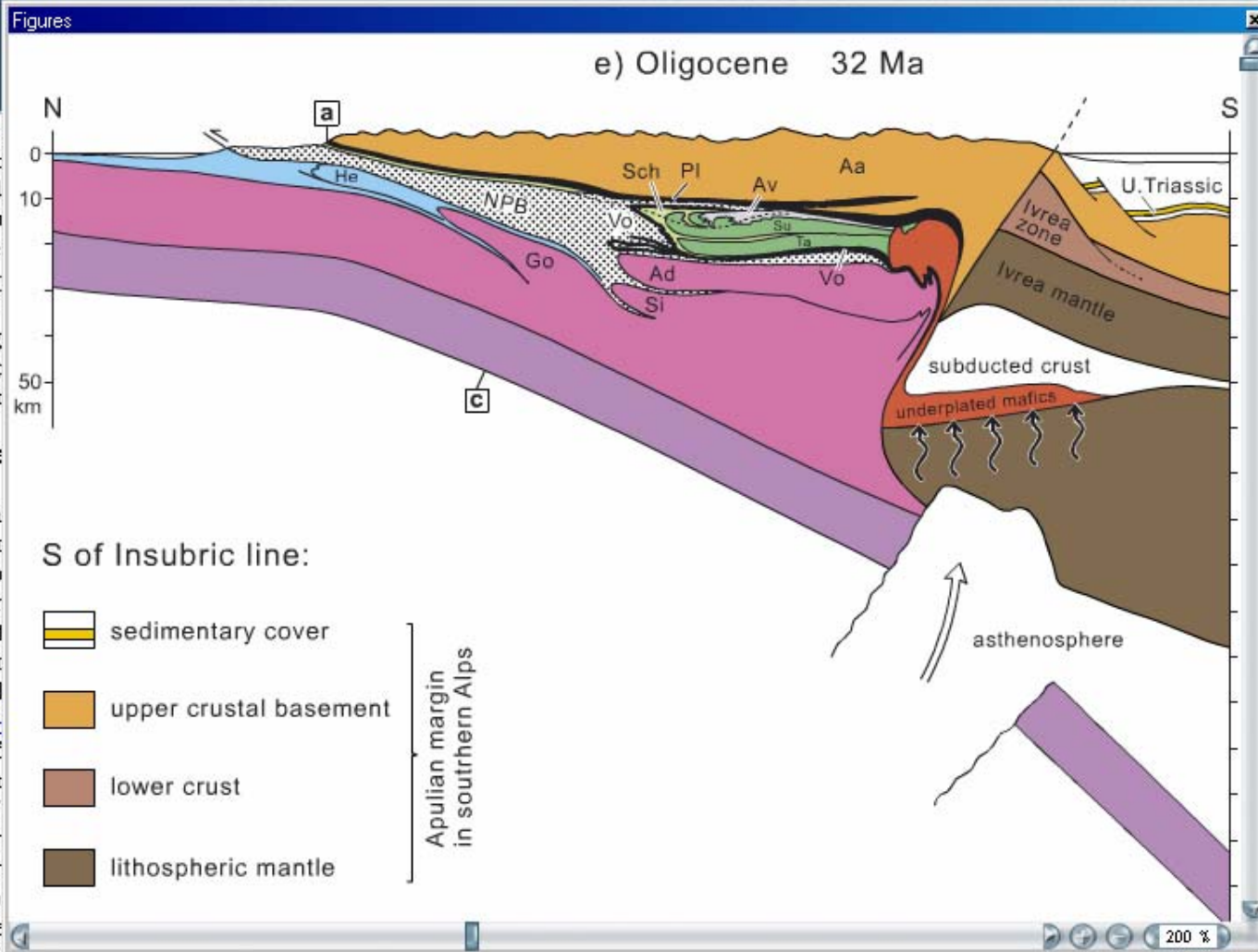
the rest of the Penninic nor into the Briançonnais margin, as documented Cretaceous-Tertiary boundary. The attribution of a pre-activity is uncertain. At the Austroalpine nappes. During the various stages ophiolites, formed a rigid part (not shown).

Early Tertiary convergence

Following closure of the which now forms an arc. Terrane entered the sub somewhat longer. After margin of Europe (the [VI.5b](#)). Penetrative deformation Briançonnais Terrane (Liguria Ocean (Fig. [IV-VI.5b](#)). Tertiary collision (50-30 Ma).

During the middle arc by the incorporation of lid formed by the Austroalpine northern foreland, 100 km N-S convergence between across the future Western convergence and collision, with west-directed movements post-dating collision (see post-collisional stage 1).

Since the Alpine nappes in Fig. [IV-VI.5a-d](#) and [5e-g](#) exclusively consist of thin upper crustal slices (basement and/or its cover) detached from their lower crustal and mantle-lithosphere, all European (and Valais) lower crust (including parts of the upper crust) must have been subducted during this collisional stage took place.



All figures cited in the text can be directly accessed via links and pop up in a separate window.

Transect IV-VI

well. In section C-C' through the 3-D lithospheric model, the Adriatic lower lithosphere is found to have been subducted to the NE and beneath the European Plate down to a depth of 270 km. Using the suture between European and Adriatic Moho as reference, the total shortening since collision amounts to some 210 km, significantly more than observed for the SE-directed subduction along transect B-B' further to the west. There is no indication for a detachment of the subducted lower lithosphere within the observed depth range. Transect C-C' reaches the Po Plain low-velocity anomaly, seen much clearer on Fig. IV-VI.12, near its southwestern end.

The 3-D tomographic model leaves no doubt that the dip direction of the subducted slab flips from a SE to a NE direction between transects B-B' and C-C', this flip occurring over a relatively short distance

of about 80 km and between the two occurs roughly beneath the Giudicaria

structure of the Alpine mantle-lithosphere the locations of these ideally chosen mantle

Fig. IV-VI.14c, except for transect A-A'. Hence, is in Fig. IV-VI.14. lithospheric features that are nearly identical in the sense of subduction and also the Fig. IV-VI.14b follows NFP 20 EAST/EGT the eastern margin of the Western Alps. It data (Holliger and Kissling 1992; Pfiffner et al. edge of relatively rigid lower Adriatic crust of European lower crust was possibly transect V. However, teleseismic tomography Fig. IV-VI.14b, the upper or lower boundary of subducted high-velocity lower

Western and Eastern Alps can be seen along a own in Fig. IV-VI.14c. Features of lower replaced by a relatively narrow collisional caria Line joins the Periadriatic Lineament are here dramatically steep and the structure as commonly seen in sediments (Dohrin and

Figures

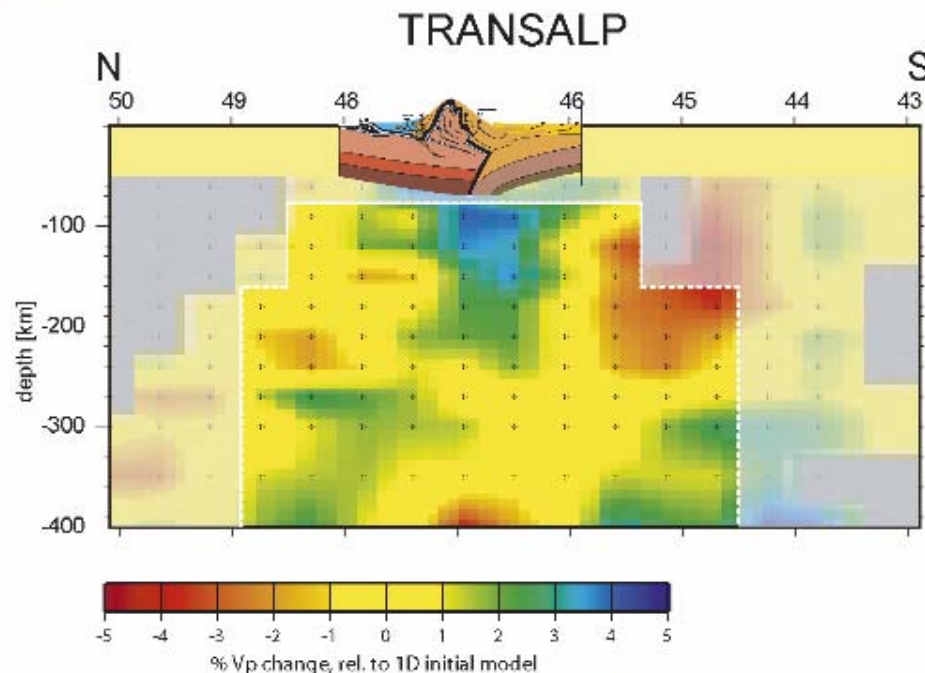


Fig. IV-VI.14: Combined crustal and lower lithosphere transects along the Alpine margin (Fig. IV-VI.3b, c, d); after Kissling et al. (in press). For location see Fig. IV-VI.1.

All figures can be zoomed and their windows scaled.

Transect IV-VI

slab rolling along the main and lateral lines, adjacent to the Tauern pop-up structure, while the alternative interpretation given by TRANSALP Working Group (2002) emphasizes displacement along a thrust at the base of the Tauern window, referred to as "Sub-Tauern ramp". In terms of the deep structure, the transect given in Fig. IV-VI.2d completely differs from that given by TRANSALP Working Group (2002). The TRANSALP section of Fig. IV-VI.2d,e -corresponding to TRANSMED Transect VI- show the Apulian Moho as descending northward under the European lithosphere, as indicated by the tomographically defined lithospheric configuration (Lippitsch 1992; Lippitsch et al. 2003; Kissling et al. in press).

This re-interpretation explains the lack of a first-order separation between Southern Alps and External Dinarides. Such a separation would be expected if the Alps and Dinarides would still exhibit opposite subduction polarities, as they did during the Eocene. Since no separation is visible between Southern Alps and Dinarides (Fig. III.11.4), both are expected to presently occupy the same lower plate. We interpret both of these first-order subduction polarities as having been initiated around 20 Ma ago and as being related to a change in subduction polarity along the Periadriatic lineament (including the Southern Alps and northern Dinarides and the Brenner Line). As discussed later, this change in subduction polarity was initiated by strike-slip movements along the lineament between the Southern Alps and northern Dinarides (Fig. VI.1) and the northern rim of the Austroalpine nappes (Apulian). This indicates that the change in subduction polarity, which only concerns the southern part of the Alpine system, the northern deformation front of the Alps can be traced back to the initiation of subduction rollback and slab break-off was initiated at about 20 Ma ago and to have occurred between the

Note that the polarity of the subduction is reversed (Fig. VI.1) and the northern rim of the Austroalpine nappes (Apulian) is in the east. This indicates that the change in subduction polarity, which only concerns the southern part of the Alpine system, the northern deformation front of the Alps can be traced back to the initiation of subduction rollback and slab break-off was initiated at about 20 Ma ago and to have occurred between the

TRANSMED Transect VI (see Fig. IV-VI.2e) best illustrates the Austroalpine nappe stack that preserved a thickness of some 10-20 km in the area east of the Tauern window, an area that lacks substantial exhumation by late stage thrusting and/or orogen-parallel extension during the Tertiary. In this transect the Koralpe-Wölz high-pressure nappe system is interpreted as representing a former extrusion wedge located between the Silvretta-Seckau nappe system in its footwall and the Ötztal-Bundschuh and Drauzug-Gurktal nappe systems in its hanging wall. This extrusion channel exhumed high-pressure units that formed during the subduction of the western embayment of the Meliata Ocean (Figs. IV-VI.2b,c) during the Late Cretaceous orogeny.

2.3 Evolution of the Alpine system and its forelands

The following discussion focuses on the timing of the subduction and exhumation of the Alpine system.

All texts can be printed either complete or just in part. [View more](#)

Transect IV-VI

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high atlas

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- ☒ Transect III ☒ Transect VII
- ☒ Transect IV ☒ Transect VIII

all transects

media type

- ☒ text
- ☐ map
- ☐ image
- ☐ reference

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 **Additional Information I.12:** the [High Atlas](#) (Ziz river valley, Jebel

 **Figure I.7:** Central [High Atlas](#), based on new in-situ

 **Figure I.7:** the [High Atlas](#) (18% at the Imilchil

 **Figure I.7:** The [High Atlas](#) is modeled as a

 **Text I:** - [High Atlas](#) - Sahara

 **Text I:** the [High Atlas](#), and finally reaches the

 **Text I:** the [High Atlas](#) and also of the

 **Text I:** the [High Atlas](#) segment and Charles

 **Text I:** the [High Atlas](#). The Panafrican crust is

 **Text I:** the [High Atlas](#) showing a folded and

 **Text I:** the [High Atlas](#) (Fig. I.2). It comprises

 **Text I:** the [High Atlas](#), the crust shows a

 **Text I:** the [High Atlas](#) (70 km according to Seber

 **Text I:** and [High Atlas](#) chains (Fig. I.2). North

 **Text I:** the [High Atlas](#), weakly oblique to the

 **Text I:** western [High Atlas](#) branch to the south

 **Text I:** central [High Atlas](#) from the Kasba-Tadla

The search offers fast and systematic access to all topics and data of the atlas.

Additional Information



Bathonian reef facies of the High Atlas (Ziz river valley, Jebel Assameur n'Aït Fergane, 100 km east of the transect)

Photo: D. Ponsard, Museum National d'Histoire Naturelle, Paris.

2° 56' W 3°

36° 28' N 36°

3° 23' W

36° 55' N

3° 02' W

37° 00' N

SSE

NNW

SSE

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00

Internal Betics

Iberian Meseta

0

nS-n6

Malaguide

T-E2

PZ-E2

Alpujarride

Nevado-Filabride

Alpujarride

Texts, references, figures and the transects themselves are searchable and can be accessed directly via link.

Alboran Domain Cr

text

legend


index map


tectonic



zoom 100 %

Sitemap


 Transect I


 Chronolithostratigraphic Transect I


 Tectonic Transect I


 Text I


 Transect II

 Transect III


 Transect IV

 Transect V

 Transect VI


 Transect VII


media type


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 Additional Information I.1


 Additional Information I.2


 Additional Information I.3


 Additional Information I.4


 Additional Information I.5


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The sitemap gives an overview about the structure and all information, texts and figures contained in the CD.

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